



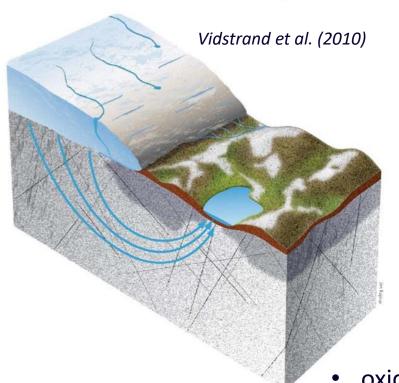
Oxygen intrusion in fractured rocks

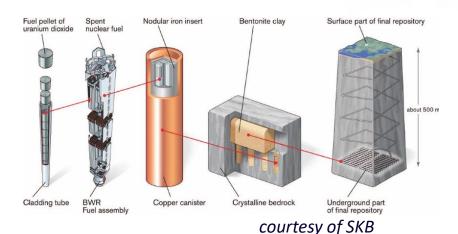
How natural analogues and numerical numerical models can work together

Jorge Molinero / NAWG 15 Workshop / Prague. May, 2017

Motivation

 The infiltration of oxygen (e.g. due to ingress of oxygenated glacial melt water) into initially anoxic fractured media is an issue of concern for the safety of nuclear waste respositories





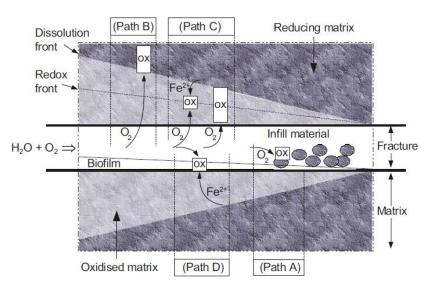
oxidizing conditions would **increase the solubility** and **mobility** of many radionuclides and could **corrode** copper canisters

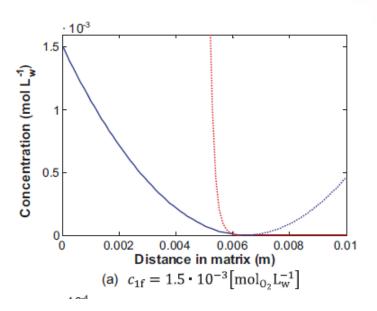


Background: Previous models

- Due to the stiffness of the underlying system of equations, existing numerical models are based on simplified 1D or 2D geometries...
- ...and on **equivalent homogeneous** parameters

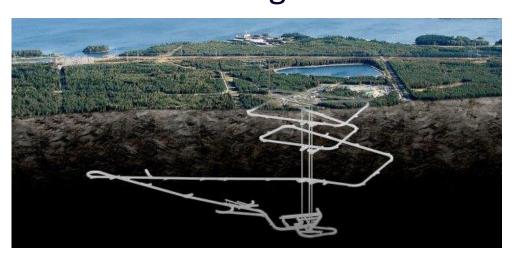
Sidborn et al. (2010)







But we are telling the site looks like this:



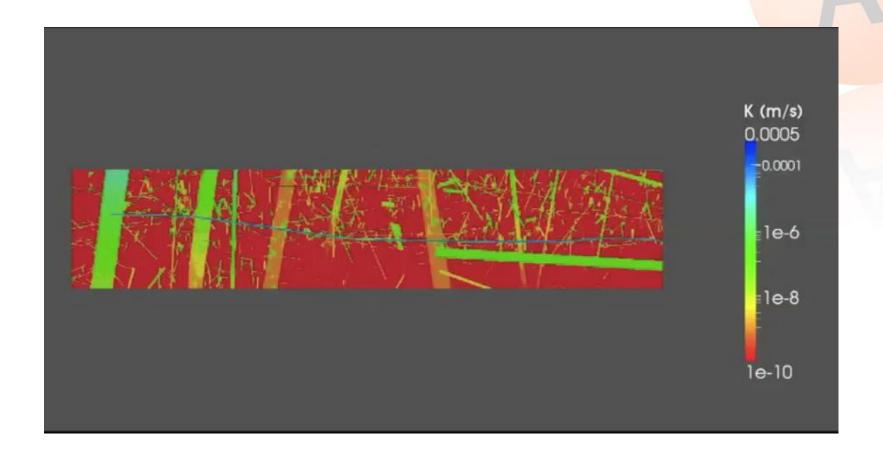








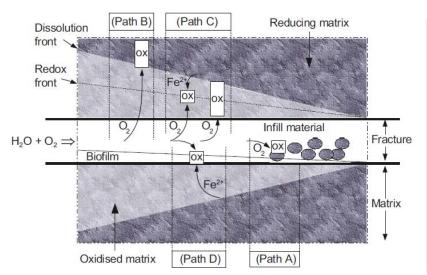
And we construct sofisticated SDM...

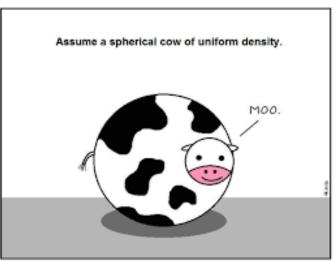


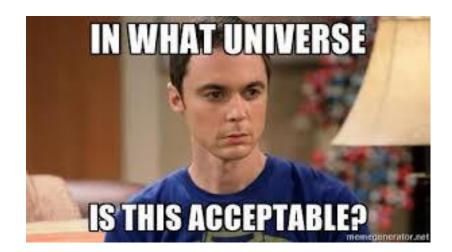


And finally we use something like this:

Sidborn et al. (2010)











We (science) know how oxygen consumption works



• **Chlorite dissolution** (Lowson et al., 2005)

$$k_{chl,diss}[mol/m_m^2 \cdot s] = 10^{-9.79} \cdot c_{H^+}^{0.49} + 10^{-13.00} + 10^{-16.79} c_{H^+}^{-0.43}$$

$$r_{Fe}[mol_{Fe}/m_b^3 \cdot s] = n_{stoich} \cdot A_s \cdot k_{chl,diss} \left(1 - \frac{c_{Fe}}{c_{Fe,sol}}\right)$$

Homogeneous oxidation of ferrous ions (Stumm & Lee, 1961)

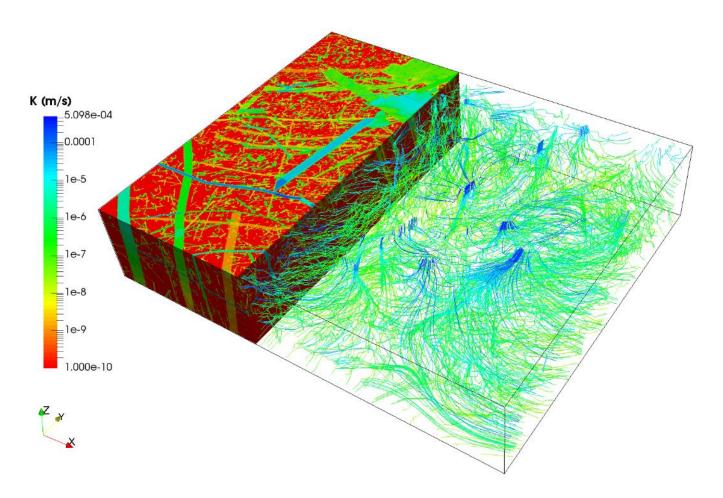
$$4Fe^{2+} + O_2 + 4H^+ \leftrightarrow 4Fe^{3+} + 2H_20$$

$$r_{ox}[mol_{Fe}/L \cdot s] = k_r \cdot c_{Fe} \cdot P_{O_2} \cdot c_{OH}^2$$



We (science) know how to represent hydrogeology in FR

A²¹

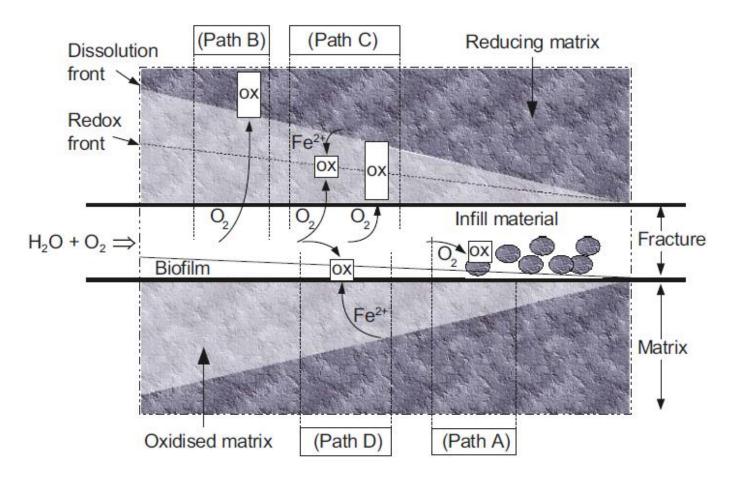


regional hydrogeological model of Forsmark



So again... why?

Sidborn et al. (2010)





It is just (or mainly) a technological barrier

Software:

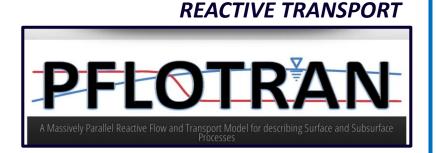
FLOW











Hardware:









Then, we incorporate our best knowledge

	Boundary water	Initial water
рН	5.8	7.3
pe	15.1	-2.0
Total concentration (mol/L)		
Al^{3+}	-	1.1×10^{-7}
Br ⁻	-	4.5×10^{-5}
HCO_3^-	2.0×10^{-6}	6.5×10^{-3}
Ca^{2+}	4.5×10^{-6}	9.5×10^{-4}
CI-	2.6×10^{-5}	3.1×10^{-2}
F ⁻	_	7.6×10^{-5}
Fe^{2+}	-	2.7×10^{-6}
K ⁺	1.0×10^{-5}	6.1×10^{-4}
Li ⁺	_	3.2×10^{-6}
Mg^{2+}	4.1×10^{-6}	2.9×10^{-3}
Na ⁺	7.4×10^{-6}	3.3×10^{-2}
SO_4^{2-}	5.2×10^{-6}	2.1×10^{-3}
SiO_2	_	1.3×10^{-4}
Sr ²⁺	_	8.1×10^{-6}
tracer	1.0×10^{0}	1.0×10^{-10}
O_2	1.5×10^{-3}	-

- 17 primary species, 29 secondary species, two reacting minerals (calcite and chlorite) and a homogeneous aqueous oxidation reaction
- 70,125,000 transport degrees of freedom



And parameterization based on SC data:

 Löfgren and Sidborn (2016) mapped minerals in 2071 open fractures and provided statistics of mineral coverage and mineral thickness for chlorite, calcite and pyrite

			All Forsmark
	Total number of	of	
	Fractures	$n_{\rm tot}$	2071
	Calcite	$n_{ m quant}$	673
		$n_{\rm curl}$	1177
	Chlorite	$n_{ m quant}$	505
$\alpha = 0.24$		n_{oual}	1077

d_{mean}	=	1.3	•	10^{-4}	m

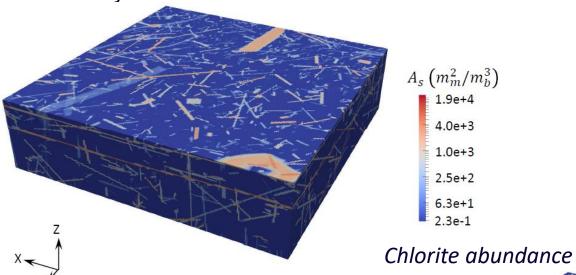
	$d_{ m mean}$	
CDF	Calcite (mm)	Chlorite (mm)
Min	0.001*	0.001*
1 %	0.001*	0.005
5 %	0.002	0.015
10 %	0.004	0.033
30 %	0.015	0.089
50 %	0.040	0.130
70 %	0.090	0.195
90 %	0.228	0.329
95 %	0.406	0.500
99 %	1.330	2.300
Max	2.000	7.000
Mean	0.107	0.216
Std	0.230	0.444



Löfgren, M., Sidborn, M., 2016. Quantitative mapping and statistical evaluation of fracture minerals in the granitic bedrock at Forsmark, Sweden. Mineralogy and Petrology, 110, 663-680.

Geochemical parameterization looks like this:





- Reactive transport calculation ran out to 340 y
- A total of about 33,000 h (3.7 years)
 of supercomputing time were
 consumed





All details of the model:





Contents lists available at ScienceDirect

Journal of Contaminant Hydrology





Continuum-based DFN-consistent numerical framework for the simulation of oxygen infiltration into fractured crystalline rocks



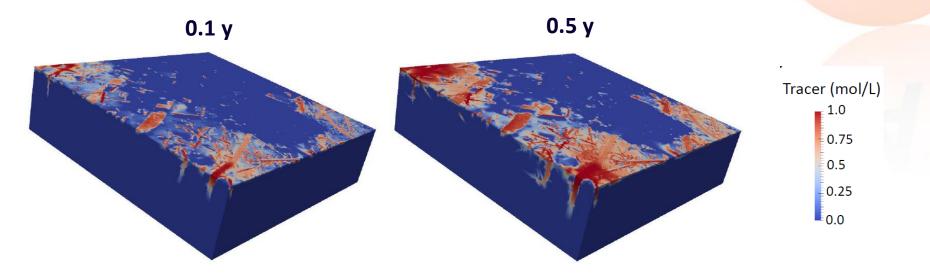
Paolo Trinchero^{a,*}, Ignasi Puigdomenech^b, Jorge Molinero^a, Hedieh Ebrahimi^a, Björn Gylling^b, Urban Svensson^c, Dirk Bosbach^d, Guido Deissmann^d

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Computed results: Conservative tracer



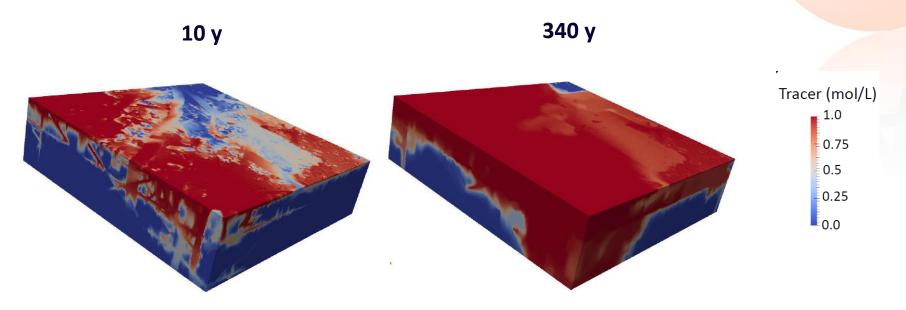


 At early times, the tracer infiltrates from the recharge areas and is channeled along transmissive fractures and deformation zones



Computed results: Conservative tracer



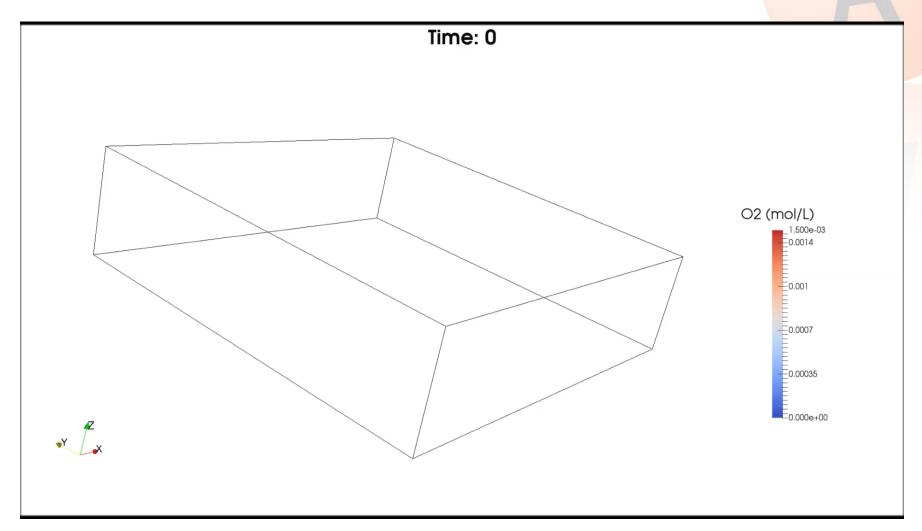


• It progressively infiltrates deep into the bedrock and reaches most of the **discharge** areas



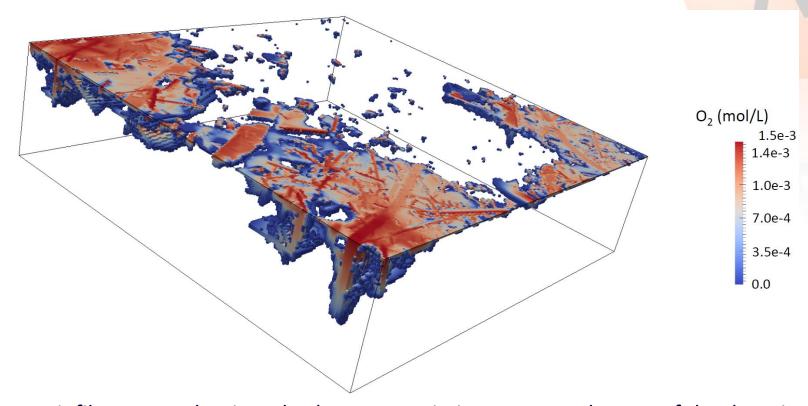
Computed results: Oxygen







Computed results: Oxygen



- Oxygen infiltrates predominantly along transmissive zones => the rest of the domain remains anoxic
- Maximum penetration depth is 90 m approximately.

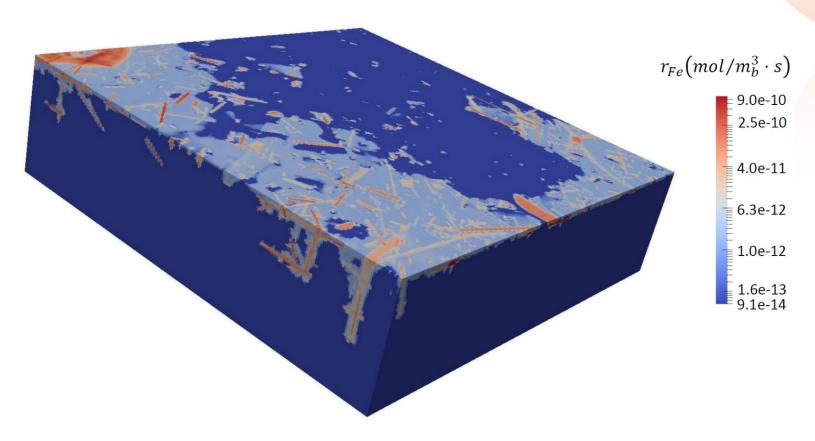


Computed results: pe & pH 15.1 14.0 10.5 7.0 3.5 0.0 -2.0 рΗ 10 9.4 8.9 8.4 7.9 7.5



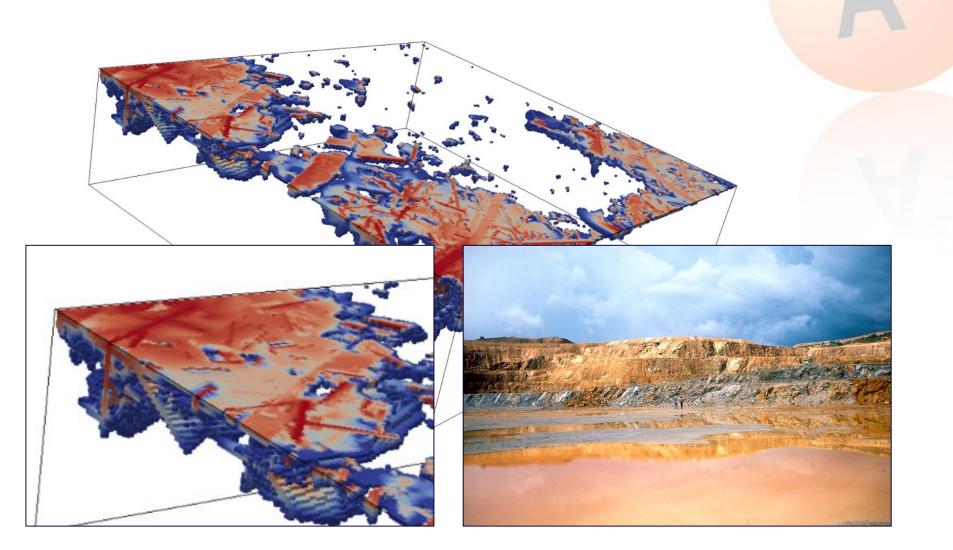
Computed results: Chlorite dissolution (rate)





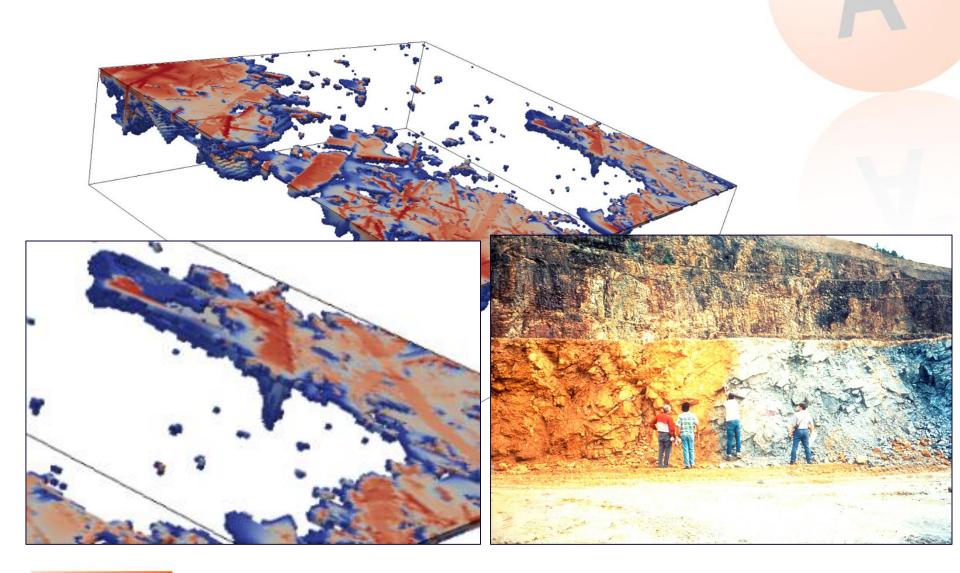


So now we have a better model... should we trust it?



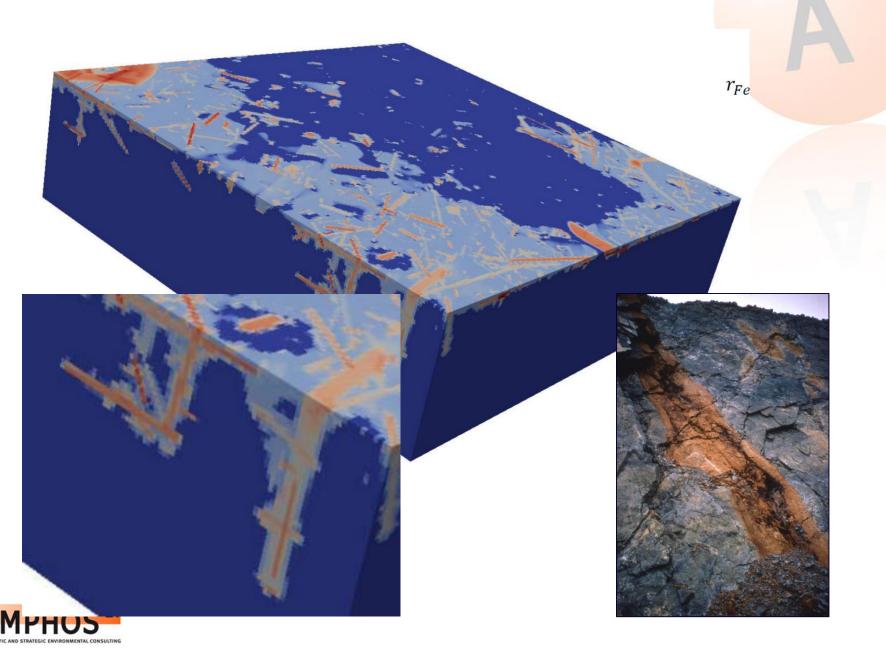


So now we have a better model... should we trust it?





So now we have a better model... should we trust it?



Conclusions

The results of the Forsmark model show that **quick penetration** of oxygen is observed only **along transmissive zones**

Steady state conditions is attained after around 10y => maximum penetration depth of **90 m**

The numerical model show results **highly consistent** with the qualitative observations of a **natural analogue** of the same processes we want to simulate (Poços de Caldas – Oxidative fronts)

Models must be adapted to the **best available technology**.

Realism and confrontation with actual **observations in nature** are the keys for building **model confidence**.



Credits

- Paolo Trinchero, Jorge Molinero & Hedieh Ebrahimi => Amphos 21
- Ignasi Puigdomenech & Björn Gylling => SKB
- Urban Svensson => CFE
- Guido Deissmann & Dirk Bosbach => Forschungszentrum Jülich

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- Thanks for your attention -

